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Original scientific paper



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Abstract

Middle Triassic deposits in the Karst Dinarides (south Lika area, near Gračac) were investigated for petrological, sedimentological and biostratigraphic determinations. Within two continuously recorded sections of sedimentary carbonates and interlayered pyroclastic rocks (Mali Kunovac and Kunovac sections) reflecting variable sedimentary-magmatic environments, valuable cephalopod fauna ranging from the middle Illyrian to the upper Illyrian age was found. The most significant was the recovery of the ammonoids *Asseretoceras* sp., indicating the upper part of the Illyrian Paraceratites trinodosus Zone, and *Reitziites reitzi* species which marks the upper part of the Illyrian *Reitziites reitzi* Zone. The discovery of the *Reitziites reitzi* species was not known from the Karst Dinarides until this finding.

Keywords:

Karst Dinarides, Croatia, Middle Triassic, biostratigraphy, ammonoids

1. Introduction

Biostratigraphic zonation according to ammonoids has a very high resolution and is still considered one of the best correlative principles for the Middle Triassic. There are many recent works that consider ammonoids as the important correlative marker in spite of the wide use of conodonts in biostratigraphic correlations. The Anisian-Ladinian boundary was defined by **Brack et al. (2005)** by means of ammonoids/conodonts. Our investigation was focused on ammonoid bearing Middle Triassic strata in the Karst Dinarides. The investigated area is located at the southern part of Lika, in Croatia. In the same area, the already known cephalopod bearing strata was recorded by **Koch (1911)** and **Salopek (1914)**. **Salopek (1914)** collected a large number of cephalopod specimens and determined the existence of *Ptychites Mojsisovics*, *Ceratites de Haan*, *Proarcestes Mojsisovics*, *Pinacoceras Mojsisovics* and *Gymnites Mojsisovics* genera. According to the specimens collected from the green clastic rocks, **Salopek (1914)** determined the investigated area is of transitional Anisian-Ladinian age. Until today, no continuous sampling of ammonoids was conducted and this study therefore presents a new and precise biostratigraphic definition of the Middle Triassic. The two sections were investigated by means of petrography, sedimentology and ammonoid biostratigraphy.

2. Geological setting

The investigated area is located in the southern part of Lika, in Croatia. In a geological sense, the area represents the central part of the Karst Dinarides. The wider area of Kunovac has been investigated in the frame of basic geological mapping, and the area of our interest covers the parts of four sheets of the Basic Geological Map 1:100 000: Udbina (**Šušnjar et al., 1973**), Drvar (**Šušnjar et al., 1978**), Obrovac (**Ivanović et al., 1973**) and Knin (**Grimani et al., 1972**). All four sheets overlap in the wider area of Kunovac Creek (see **Figure 1**). According to the Geology of Knin sheet (**Grimani et al., 1975**), the investigated area is composed of two chronostratigraphic units (T_2^1 – Anisian and T_2^2 – Ladinian, comprising various lithologies: (1) Anisian crystalline and white-brown limestones and grey-white poorly layered dolomites (T_2^1), up to 550 m thick; and (2) Ladinian white-grey to white crystalline limestones and coarse-crystalline dolomites interlayered with volcanogenic sediments (T_2^2), up to 750 m thick. Towards the north, Lower Triassic units appear: T_1^2 – clayey and sandy limestones, marls and dolomites („*Campiler beds*“), up to 470 m thick; and T_1^1 – micaceous and silty sandstones, various types of clayey and sandy limestones and marls („*Seisser beds*“), up to 500 m thick. Towards the south, Upper Triassic dolomites, as well Lower Jurassic limestones and dolomites, follow. The investigated sections are located at the structural unit comprising of conformable succession, ranging from the early Triassic to the late Jurassic. The strata are

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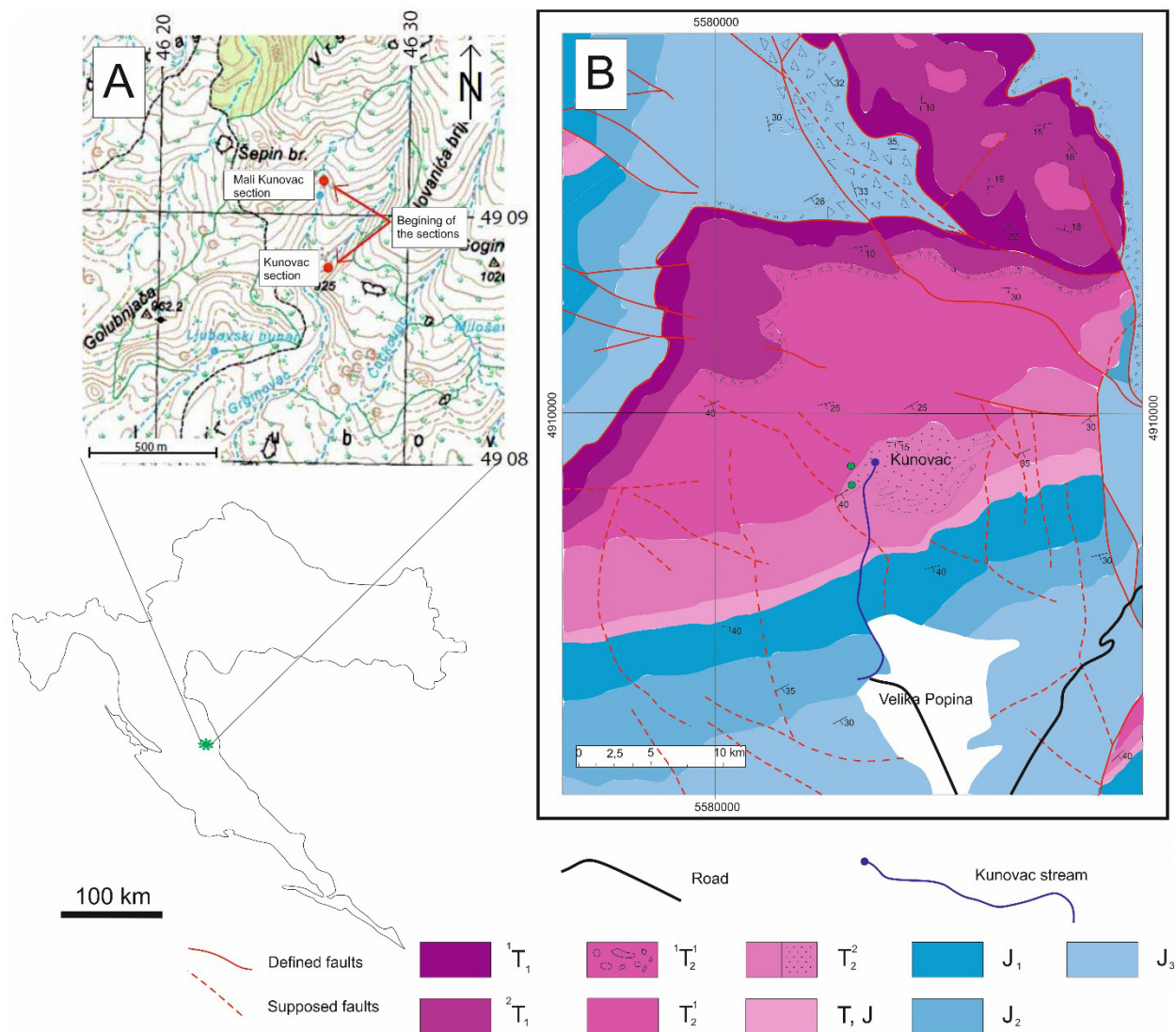


Figure 1. A map of the investigated area. An outline of the Republic of Croatia is presented in the lower part of the figure, with a green complex star marking the investigated area. A) A topographic map of the investigated area, with red circles marking the beginnings of the recorded sections. B) A geological map of the investigated area, combined from contiguous/marginal parts of the sheets of the Basic Geological Maps of SFRY 1:100 000: Udbina (Šušnjar et al., 1973), Drvar (Šušnjar et al., 1978), Obrovac (Ivanović et al., 1973) and Knin (Grimani et al., 1972). Green circles mark the investigated area and the localities of the recorded sections.

dipping to the south-southeast. In the north and north-west, the structural unit is bordered by a fault, bringing the early Triassic strata in contact with the late Jurassic deposits. The two recorded sections, named Mali Kunovac and Kunovac, were recorded in creek gullies and they can be correlated into one continuous section, with a covered interval in between. The older section is represented by the deposits of Mali Kunovac with its beginning marked on **Figure 1**.

3. Methods

Two sedimentary sections, named Mali Kunovac and Kunovac, were measured and sampled during detailed field work. Samples of different rock types were collected for micropetrographic and facies analysis. The

thin sections were prepared at the Department of Mineralogy, Petrology and Mineral Resources of the Faculty of Mining, Geology and Petroleum Engineering, University of Zagreb, by the standard procedure, and stained with Alizarine red-S and K-ferricyanide to distinguish calcite and dolomite minerals. Twenty-two thin sections were prepared and analysed using a Leica DM/LSP polarizing microscope to determine micropetrographic features of the sampled rock types. Samples of different cephalopods were collected for biostratigraphic and paleoecologic investigations. Samples were prepared and determined at the Paleontological Department of the Croatian Natural History Museum. The genus and species of different cephalopods were determined according to the specific whorl morphological characteristics, including ornamentation and lobe line type.

4. Results

4.1. Investigated sections: Mali Kunovac and Kunovac

The two sections, Mali Kunovac and Kunovac, represent a conformable Anisian/Ladinian transitional sedimentary succession (see **Figures 1 and 2**). The two sections are divided by an approximately 10 meter thick covered interval. Since no tectonic breaks are recorded and the layer orientation of both sections is the same (the general dip is towards the southeast, with inclinations of $\sim 25^\circ$), they are considered as a continuous depositional sequence, with the Mali Kunovac section representing the older, and the Kunovac section representing the younger strata (see **Figure 2**).

At the very base of the Mali Kunovac section, massive grey limestones occur (see **Figure 2**). Layers range from 1 to 2 meters in thickness (see **Figure 3A**). They are determined as bioclastic floatstone to grainstone. Fossil content is dominated by bivalves, brachiopods, crinoids, foraminifera, corals, echinoids, green algae, gastropods and bryozoans (see **Figure 3B and 3C**). Sometimes, even microbial (possibly cyanobacterial) fabric can be distinguished. A subordinate amount of various sized intraclasts are recognized. Lime mud-rich varieties show wackestone/floatstone fabric whereas sparite-rich varieties are dominated by grainstone fabric and sparry calcite cement.

The massive limestones are conformably overlain by a 2.5-meter-thick red nodular limestone bed (see **Figure 3D**). The interval is intensively silicified and contains pyroclastic detritus. A micropetrographical analysis revealed that the lowermost part of this interval consists of bioclastic grainstone with up to 30% of pyroclasts. The limestone component is presented by recrystallized intraclasts, bivalve fragments, crinoids and radiolarians. Pyroclastic material dominantly consists of hyaline lithic fragments, and subordinately of feldspar and quartz crystalloclasts. In the upper part of the red limestone interval, biomicritic limestone with dominantly pelagic fauna (thin-shelled bivalves, radiolarians and crinoid fragments), is recognized.

Red limestone is overlain by a meter-thick interval, consisting of the interlayered limestone and pyroclastic beds. Sometimes, horizontal lamination can be observed. Limestones from this interval are determined as biomicritic and lithoclastic limestones containing various amount of pyroclastic material. Biomicritic interlayers/laminae contain radiolarians, thin-shelled bivalves and crinoids, supported by lime mud (see **Figure 4B**). Lithoclastic limestones contain irregularly shaped limestone lithoclasts. Lithoclasts are dominantly of micritic and biomicritic types. Between the lithoclasts, sparry calcite cement was precipitated (see **Figure 4A**). Sometimes, a thin lamination with pyroclastic material occurs between the limestone laminae. Limestones are silicified and re-

crystallized to a various extent. Pyroclastic material in the limestones are feldspar and quartz crystalloclasts, as well as hyaline lithoclasts. The majority of the ammonoid specimens were collected in this interval.

The top of the Mali Kunovac section is marked by a 4.5-meter-thick pyroclastic interval. The pyroclastic interval consists of greenish tuff layers (locally called the *pietra verde*) composed of fine ash material, and brown layers containing material of coarse ash size (see **Figure 4C**). Fine ash tuff layers have homogenous textures, and they are composed of very fine ash size glass shards (see **Figure 4D**). Exceptionally rare feldspar crystalloclasts were recognized. Coarse ash tuff, determined as crystalloclastic tuff, consists of plagioclase and quartz crystalloclasts, and subordinately of pumice fragments and lithoclasts. Most of the plagioclase crystalloclasts are altered to calcite and prehnite to a various extent. Irregularly shaped quartz crystalloclasts often exhibit oval cav-

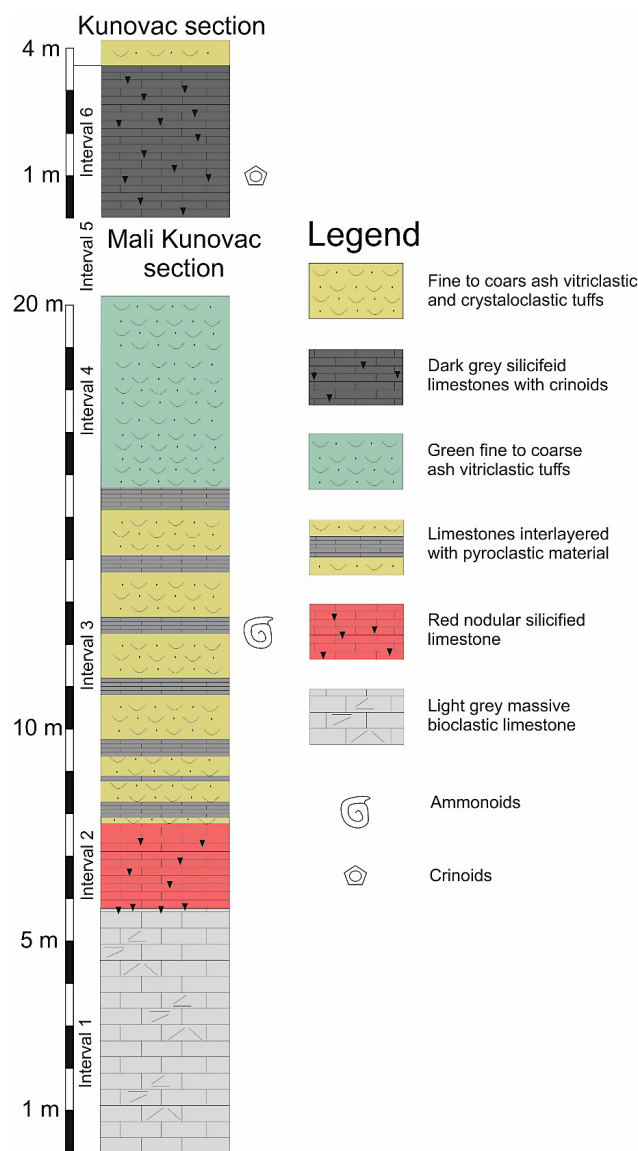


Figure 2. Recorded sections in the investigated area with varied lithology and determined macrofossils.

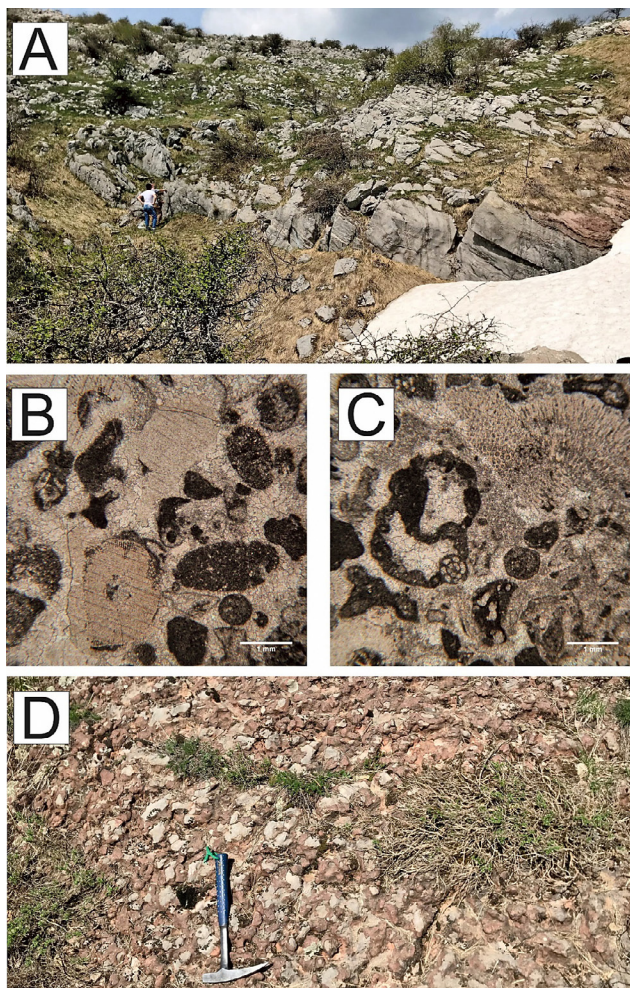


Figure 3. Field photographs and photomicrographs of the lithologies recognized in the Mali Kunovac section.

A) A field photograph of the massive grey limestones at the beginning of the Mali Kunovac section.

B) and **C)** Photomicrographs of grainstone/floatstone limestone types containing dominantly shallow marine fossils of foraminifera, green algae, corals and echinoids.

D) A field photograph of the red limestone layer with a nodular texture.

ities. Lithoclasts are dominantly of andesitic composition. Surprisingly, crinoid fragments are sometimes present. Between the pyroclasts, a matrix composed of very fine greenish hyaline material, can be observed.

The covered interval between the two recorded sections is probably composed of severely altered fine ash tuffs covered with vegetation. Fragments of fine greenish tuffs were found all along the covered interval.

The oldest rocks of the 3.6 m thick Kunovac section are platy limestones, rich in crinoid fragments (see **Figure 2**). Beds are up to 0.5 m thick, with an irregular or nodular silicification pattern observed within the layers. The silicified parts are yellowish to brownish in colour. The limestones in the Kunovac section are determined as lithoclastic limestone in the lower part, and packstone with pelagic fossils in the upper part of the interval.

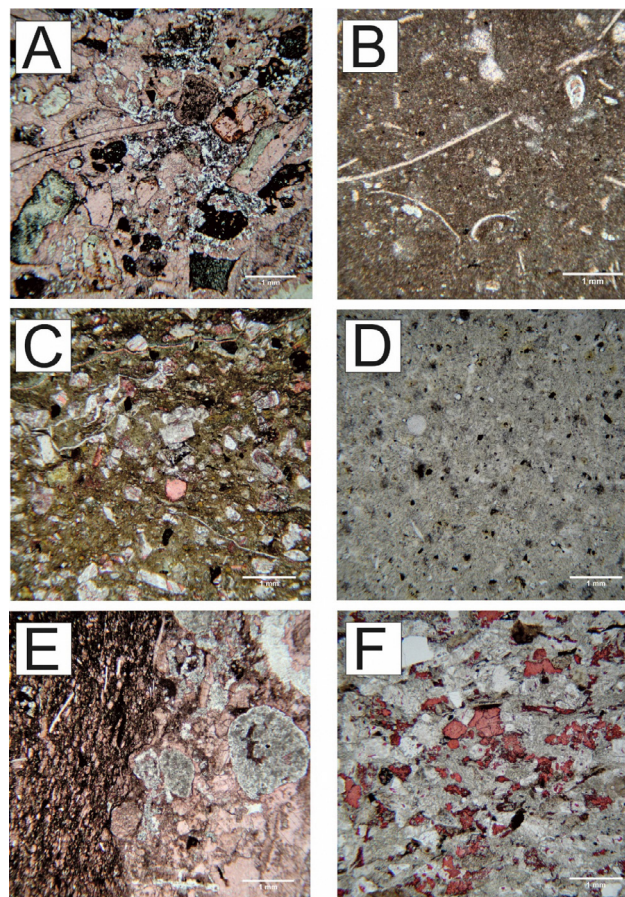


Figure 4. Photomicrographs of the samples from the Mali Kunovac section (**A**, **B**, **C** and **D**) and the Kunovac section (**E** and **F**). **A)** Lithoclastic limestone from the Mali Kunovac section containing lithoclasts of shallow marine biomictite and different pyroclastic particles (white). **B)** Biomictic limestone – wackestone/packstone with pelagic fauna (radiolarians, thin-shelled bivalves). **C)** A sample of the pyroclastic layer found in the Mali Kunovac section. Note the greenish matrix composed of fine hyaline fragments. **D)** Fine ash tuff composed of very fine glass shards – Mali Kunovac section. **E)** Recrystallized limestone (right) in contact with pyroclastic material (left) from the Kunovac section. **F)** Crystalloclastic tuff from the Kunovac section containing feldspar and quartz crystalloclasts.

Lithoclastic limestone, found in the lower part of the Kunovac section, contain extremely unsorted lithoclasts of various shallow water limestones with preserved marine fauna (echinoids, green algae, foraminifera, bivalves). Lithoclastic limestone beds are intensively silicified and recrystallized (see **Figure 4E**). Lithoclastic limestone is overlain by autochthonous limestones containing pelagic fossils and pyroclastic material. Limestones found in the upper part of the interval are determined as biomictite, with a pronounced nodular silicification pattern. Fossils of thin-shelled bivalves, radiolarians, sponge spicules and crinoids were noticed. Pyroclastic material consists of plagioclase and irregularly shaped quartz crystalloclasts. Quartz crystalloclasts are usually finer than plagioclase. Sometimes, faint ir-

regular lamination can be observed due to a change of pyroclastic components in each lamina.

Above the recorded Kunovac section, another horizon of pyroclastics was found and determined as fine ash vitriclastic tuff and coarse ash crystalloclastic tuff. Vitriclastic tuff is green, with a homogenous texture. They are composed of very fine needle-shaped glass shards and exceptionally rare plagioclase crystalloclasts. Crystalloclastic tuffs are dominantly composed of well sorted euhedral/subhedral plagioclase crystals, and subordinately quartz crystalloclasts (see **Figure 4F**). Quartz crystalloclasts are mostly irregular, rounded, and usually exhibit oval cavities. Scarce pumice fragments and andesite lithoclasts can also be recognized. A certain amount of crinoid fragments is present as well.

4.2. Biostratigraphy

During fieldwork, several samples of cephalopod fauna were collected. Ammonoid specimens were mainly concentrated in the interval of the limestones interlayered with pyroclastics from the Mali Kunovac section, and just below the green vitriclastic tuffs of the *pietra verde* type. In the Mali Kunovac section, from the unit of the limestones with pyroclastics (see **Figure 2**) a significant number of *Ptychites* specimens were collected, as well as several specimens of the genus *Proarcestes* Mo-

jsisovics and *Epigymnites ecki*. The most significant findings in this interval are the specimens determined as the *Reitziites reitzi* ammonoid species. Based on the occurrence of the *Reitziites reitzi*, the stratigraphic position of the investigated strata is determined as the upper Illyrian Reitziites Zone (**Brack and Rieber, 1993; Mietto et al., 2003; Balini et al. 2010**). In the green vitriclastic tuffs, specimens of *Asseretoceras* sp. were discovered. According to the discovery of *Asseretoceras* sp., deposits in the investigated section are attributed to the upper part of the Paraceratites trinodosus Zone and the Reitziites reitzi Subzone (the upper part of the Reitziites reitzi Zone) (see **Figure 5**). The discovered cephalopod fauna is presented in **Figures 6 and 7**.

5. Sedimentary model

Massive grey limestones (at the base of the Mali Kunovac section – interval 1 in **Figure 2**) with the abundance of green algae, bivalves, brachiopods, corals, echinoids, gastropods and bryozoans in the beginning of the section indicate deposition in a shallow marine environment, probably in the shallow subtidal zone, since green algae are the dominant constituents. Reef organisms, such as corals and bryozoans, preserved as fragments, suggest redeposition of the fragmented debris

| Stage | Substage | Balaton Highland | | Eastern Lombardy | | |
|---------------------------------|---------------------------------|----------------------------------|---------------------------------|----------------------------------|---------------------------------|---------------------------|
| | | Zone | Subzone | Zone | Subzone | |
| Ladinian | Fassanian | Eoprotrachyceras curionii | Eoprotrachyceras curionii | Eoprotrachyceras curionii | Eoprotrachyceras curionii | |
| | Anisian | Illyrian | Nevadites secedensis | Nevadites secedensis | Nevadites secedensis | Nevadites secedensis |
| Ticinites crassus | | | | Ticinites crassus | | |
| Reitziites reitzi | | Illyrian | Reitziites reitzi | Aplococeras avisianum | Reitziites reitzi | Aplococeras avisianum |
| | | | | Reitziites reitzi | | Reitziites reitzi |
| | | | | Hyparpadites lieopoldti | | Hyparpadites bagolinensis |
| | | | | Kellnerites felsoeoersensis | | "Kellnerites" |
| Paraceratites trinodosus | | Illyrian | Paraceratites trinodosus | Lardaroceras pseudohungaricum | Paraceratites trinodosus | "Lardaroceras" |
| | | | | Asseretoceras camunum | | "Asseretoceras" |
| | | | | Paraceratites trinodosus | | Paraceratites trinodosus |
| | | | | Schreyerites ? binodosus | | Schreyerites abichi |
| Pelsonian | Balatonites balatonicus | Balatonites balatonicus | Bulogites zoldianus | Balatonites balatonicus | Bulogites zoldianus | |
| | | | Beyrichites cadoricus | | Balatonites balatonicus | |
| | | | Balatonites balatonicus | | | Balatonites balatonicus |
| Bithynian | Aghdarbandites ismidicus | Balatonites ottonis | Aghdarbandites ismidicus | Balatonites ottonis | | |

Figure 5. Biostratigraphic subdivision of the Middle Triassic, after **Vörös (2014)**. Zones marked in blue represent the biostratigraphic position of the fauna recovered in this investigation.

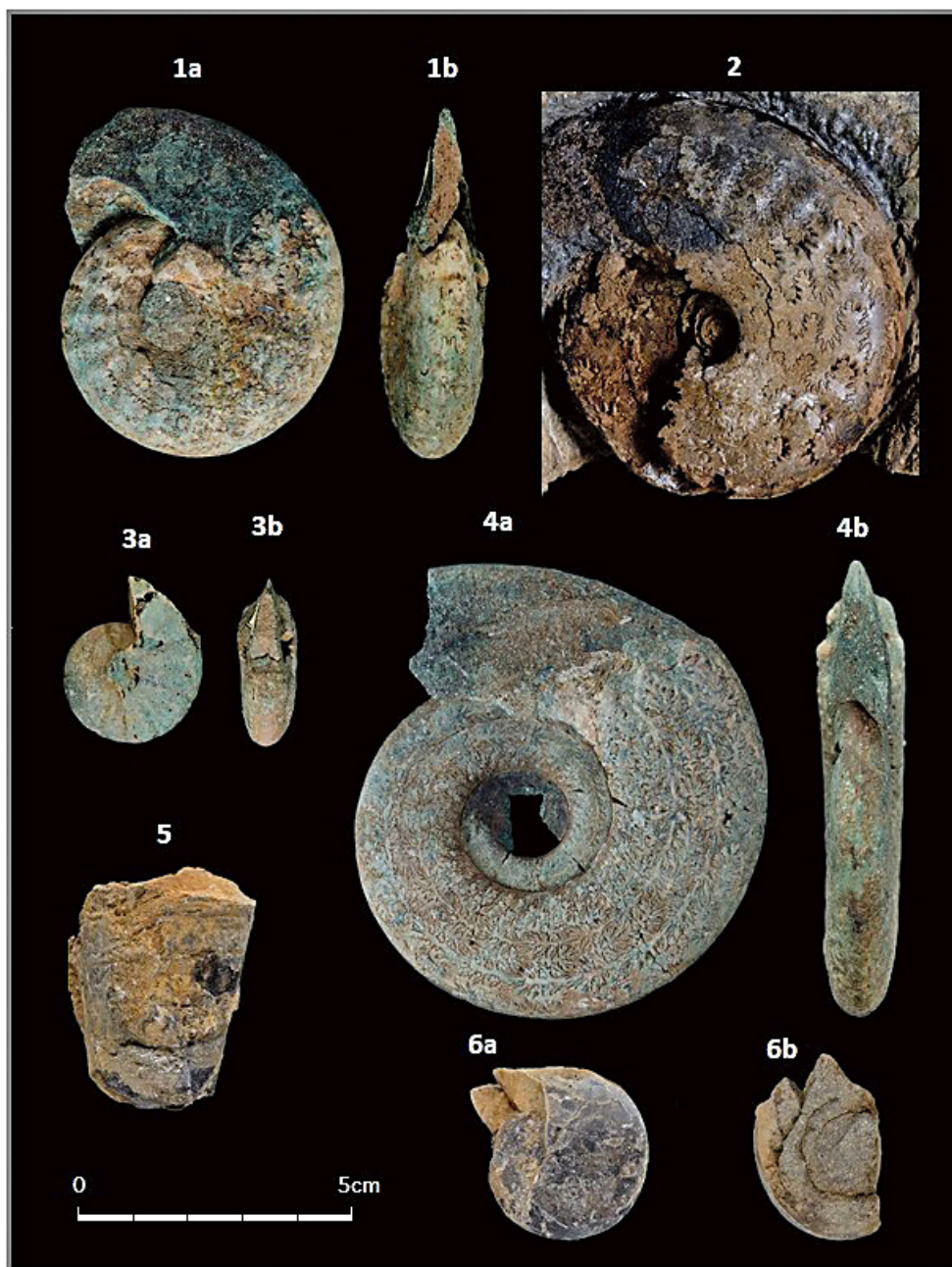


Figure 6. 1. *Ptychites oppeli* MOJSISOVICS, 1882— INV 1001, Kunovac l., Reitzi Zone (?), a: lateral view, b: ventral view, 2. *Ptychites oppeli* MOJSISOVICS, 1882— INV 1025, Kunovac, Reitzi Zone (?), lateral view, 3. *Ptychites* sp.— INV 1022, Kunovac l., Reitzi Zone (?), a: lateral view, b: ventral view, 4. *Epigymnites ecki* (MOJSISOVICS, 1882) — INV 569, Kunovac l., Reitzi Zone (?), a: lateral view, b: ventral view, 5. *Trematoceras cf. campanile* (MOJSISOVICS, 1882) — INV 1034, Kunovac l., Reitzi Zone (?), a: lateral view, 6. *Proarcestes* sp. — INV 1016, Kunovac l., Reitzi Zone (?), a: lateral view, b: cross-section.

from the nearby existing reef areas. Fragmentation was possibly caused by severe waves that also caused the re-sedimentation of the material. The deposition of carbonates in the shallow marine area was interrupted by episodes of explosive volcanism, indicated by the presence of limestones interlayered with pyroclastic deposits, just on top of the massive grey limestones (see **Figure 2**). The influence of volcanism on the Middle Triassic depo-

sition is commonly witnessed in the Dinarides (**Šćavničar et al., 1984; Kolar-Jurkovšek, 1990; Petek, 1998; Smirčić et al., 2016; Smirčić et al., 2018**) as well as in the surrounding area of the western Tethyan realm (**Brandner and Bechstädt, 1970; Obenholzner, 1991; Dimitrijević, 1997; Velledits, 2004, 2006; Kovács et al., 2011; Velledits et al., 2011**). The limestone layer at the top of the massive grey limestones with shallow wa-

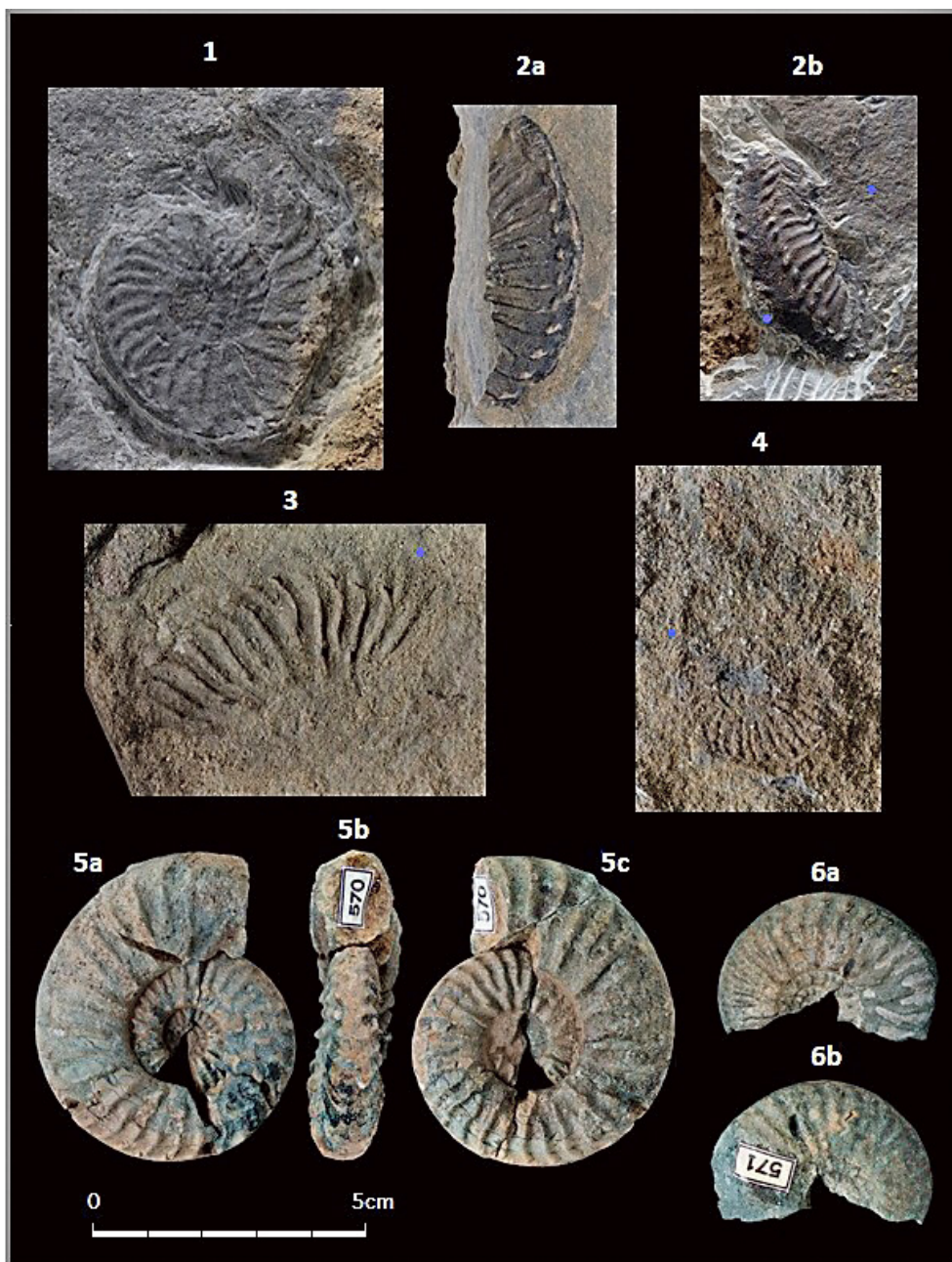


Figure 7. 1. *Reitziites reitzi* (BÖCKH, 1872) — INV 1027, Kunovac I., Reitzi Zone, lateral view, 2. *Reitziites reitzi* (BÖCKH, 1872) — INV 1029, Kunovac I., Reitzi Zone, a: lateral view, b: ventrolateral view 3. and 4. Resembling *Reitziites cf. reitzi* (BÖCKH, 1872) (as suggested by A. Vörös) — INV 1027, Kunovac I., Reitzi Zone, lateral view 5. *Asseretoceras* sp. (BALINI, 1992) — INV 570, Kunovac I., upper part of the *Paraceratites trinodosus* Zone a: lateral view, b: ventral view, c: lateral view. 6. *Asseretoceras* sp. (BALINI, 1992) — INV 571, Kunovac I., upper part of the *Paraceratites trinodosus* Zone (?), a and b: lateral views.

ter depositional characteristics contains hipidiomorphic to idiomorphic plagioclase crystalclasts and volcanic lithoclasts (interval 2 – **Figure 2**). The shapes of the crystalclasts suggest the vicinity of the primary eruption and possible air fall into a shallow marine depositional area. The volcanic activity was accompanied by tectonic movements causing the uplifting and subsiding of the blocks. The tectonic movements were probably

very intensive, and the large scaled fault movements caused drastic changes in the depositional areas. Limestones deposited in the upper part of the interval of limestones interlayered with pyroclastics contain pelagic biota supported by lime mud, indicate slow and steady suspension settling in the pelagic conditions. Drowning of the shallow marine environments, previously dominated by limestone deposition, followed. Limestones in

interval 3 (see **Figure 2**) are interlayered with crystalloclastic and vitriclastic tuffs, indicating eruptional episodes of the nearby volcanic systems (intervals 3 and 4 – **Figure 2**). The beginning of the Kunovac section, divided from the Mali Kunovac by cca 10 m covered interval (interval 5 – **Figure 2**), is marked by the lithoclastic limestones containing fragments of the shallow marine limestones. This indicates erosion of the material from the shallow marine environment and redeposition by gravity flows to the deeper marine areas (interval 6 – **Figure 2**). The erosion of the limestones was related to the tectonic activity that persisted throughout the period of deposition. The pelagic sedimentation continued after this by the suspension settling of lime mud and the dominance of pelagic fossils. The pyroclastics on top of the recorded section (interval 7 – **Figure 2**) witness the new eruptional episode. This pyroclastic interval is composed of vitriclastic and crystalloclastic tuff layers, suggesting resedimentation by gravity flows or pyroclastic density currents (**Di Capua and Gropelli, 2016**) from the shallower areas where the pyroclastic material was primarily deposited.

6. Discussion

Similar depositional features as described for the investigated sections Mali Kunovac and Kunovac, have already been documented in different parts of the Dinarides, which represents a vast marine area with small horsts (highs with shallow water, platform-like deposition) and intraplatform depressions – grabens (with pelagic-type deposition) (e.g. **Šćavničar et al., 1984; Belak, 2000; Smirčić et al., 2018**). In the investigation of the Middle Triassic deposits of the Svilaja Mts. near Sinj (cca 150 km from sections investigated in this work), **Šćavničar et al. (1984)** recognized two basalt occurrences and associated tuff horizons. The authors stated that the first basalt effusion had already started in Anisian, while the second basalt and tuff horizon belong to the Ladinian. These two mentioned volcanic/tuff horizons suggest similar time intervals as in the investigated sections described in this study. The discovery of the *Asseretoceras* sp. indicates that the lower part of the section belongs to the topmost part of the Illyrian Paraceratites Trinodosus Zone. According to the *Reitziites reitzi* species, found in the first tuff interval/interlayer from the Mali Kunovac section, the age of this short interval is the late Illyrian (*Reitziites reitzi* Zone after **Vörös, 2014; 2019**). Consequently, the volcanism in the Dinarides began already in the Anisian, as concluded also in **Smirčić et al. (2018)**. The second tuff interval, on top of the Kunovac section, certainly represents younger deposits. There is no biostratigraphic proof of its Ladinian age, so far.

Tectonic and volcanic activity in the Middle Triassic was related to the opening of the Tethyan Ocean and the disintegration of Pangea. Evidence of the volcanic activity can be traced in the whole area of western Tethys, with

its peak and termination in the late Illyrian (**Vörös, 2014; Smirčić et al., 2018**). Tectonic activity caused the intensive differentiation of the sedimentary environments forming deep marine, shallow marine and continental sedimentary environments in a relatively short distance one from another. This differentiation of the marine areas probably caused the more intensive fluctuation of the nutritive rich sea water, thus causing blossoming of the marine organisms. This also affected ammonoids since many genera appear in the late Anisian *Reitziites reitzi* Zone (*Hungarites, Parakellnerites, Aplococeras, Latemarites*), which were ecologically related to platform or peri-platform areas (**Vörös, 2014; 2019**).

Grimani et al. (1975) and **Sokač et al. (1976)** indicate that the deposits in the area of Kunovac are Ladinian in age. This study, according to the new stratigraphic division of the Middle Triassic by **Brack et al. (2005)**, proves differently. The index fossil *Reitziites reitzi* precisely indicates the *Reitziites reitzi* Zone (see **Figure 4**) (**Brack and Rieber, 1993; Mietto et al., 2003; Brack et al., 2005; Balini et al., 2010**) of the upper Illyrian and correlates well with the volcanic deposits in the other areas in the Dinarides (**Smirčić et al., 2018**).

7. Conclusion

The two recorded sections in southern Lika form a rather short but unique and significant depositional sequence. In the whole sequence, a severe and abrupt change of lithology occurs, from shallow marine limestones with abundant green algae interlayered with pyroclastics, to the pelagic limestones with thin-shell bivalves and radiolarians, accompanied with tuffs. These differences in the composition can indicate tectonic activity accompanied with volcanic eruptions at the time of the deposition. The rifting tectonic movements caused severe changes of depositional environments and an abrupt differentiation of lithological characteristics. Pyroclastic material, found in the rocks from both environments (shallow and deeper marine), indicates persistence of the volcanic activity, starting already in the Anisian (late Illyrian). Two main tuff horizons are supposed, one in the Mali Kunovac section, determined as late Illyrian in age, and the second at the top of the Kunovac section, representing younger strata. The lower part of the section is marked by the *Asseretoceras* sp., indicating the upper part of the Illyrian Paraceratites trinodosus Zone. The first tuff interval and the first finding of the *Reitziites reitzi* ammonoid species marks the upper Illyrian *Reitziites reitzi* Zone. The first record of the *Reitziites reitzi* in the Karst Dinarides directly constrains the beginning of the volcanic activity to the upper Illyrian. The occurrence of rich cephalopod fauna indicates that the deposition took place in an open marine environment with shallow marine areas existing nearby, as suggested by the presence of the shallow marine limestone lithoclasts in the pelagic environment.

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SAŽETAK

Prvi nalaz *Reitziites reitzi* markera gornjoilirike amonitne podzone u krškim Dinaridima

Srednjotrijaske naslage u krškim Dinaridima (okolica Gračaca u južnoj Lici) istraživane su s ciljem njihove petrološke i sedimentološke odredbe te odredbe njihove biostratigrafske pripadnosti. Unutar dvaju snimljenih kontinuiranih sljedova koji se sastoje od izmjene karbonatnih sedimentnih stijena i piroklastita (Mali Kunovac i Kunovac), a koji odražavaju varijabilne sedimentacijsko-magmatske okoliše njihova postanka, prikupljena je relativno malobrojna, ali vrijedna kolekcija cefalopoda srednjolirike i gornjoilirike starosti. Najvažniji su nalazi amonitne vrste *Asseretoceras* sp. koja upućuje na gornji dio ilirike *Paraceratites trinodosus* zone te vrste *Reitziites reitzi* koja označava gornjoiliriku Reitziites zonu. Pronalazak vrste *Reitziites reitzi* do sada nije bio zabilježen na području krških Dinarida.

Ključne riječi:

krški Dinaridi, Hrvatska, srednji trijas, biostratigrafija, amoniti

Authors contribution

Duje Smirčić orcid.org/0000-0002-6360-8203 (PhD, senior researcher, sedimentology/volcaniclastic rocks) along with **Nikolina Gaberšek** (MS), **Dunja Aljinović** orcid.org/0000-0002-4698-1727 (PhD, Full Professor, sedimentology/carbonates) and **Uroš Barudžija** orcid.org/0000-0002-1617-9362 (PhD, Assistant Professor, sedimentology/diagenesis) conducted fieldwork, micropetrographically analysed the thin sections, determined the lithological characteristics of the rocks and proposed sedimentary processes and a model for the investigated deposits. **Dražen Japundžić** (MSc, senior curator, palaeontology) conducted field work and made paleontological and biostratigraphic determinations of the ammonoid specimens. **Nediljka Prlj-Šimić** (MSc, senior curator, cephalopod palaeontology), **Katarina Krizmanić** (MSc, senior curator, palaeontology) and **Ivor Pavić** (MS) conducted field work investigations.